



Nanoparticulate Nickel-Hosting **Phases in Sulfidic Environments: Effects of Ferrous Iron and Bacterial Presence on Mineral Formation** Mechanism and Solid-Phase **Nickel Distribution**

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The precipitation of nickel with sulfide is an important process governing the bioavailability of Ni in natural waters, and this process has the potential to effectively remove aqueous Ni contaminants in near-surface environments. In this study, we use experimental approaches to investigate the diversity of Ni-hosting phases precipitated in sulfidic environments across a range of aqueous Ni-to-Fe ratios ([Ni]ao/[Fe]ao) and in the presence or absence of the sulfate-reducing bacteria (SRB), Desulfovibrio vulgaris. In the absence of Fe(II), the initial precipitates in abiotic experiments are found to consist primarily of polyphasic Ni-sulfides (average sizes <20 nm) with millerite (trigonal NiS) cores and amorphous shells. The precipitates' crystallinity is enhanced noticeably over a period of ~ 6 days, forming larger-sized hexagonal α -NiS, and observations of defects such as twinning and stacking faults implicate a formation pathway via reassembly of fine nanoparticulate precursors. By comparison, in the presence of SRB and in the absence of Fe, more crystalline phases such as polydymite (Ni₃S₄) and vaesite (NiS₂) are also precipitated in addition to the monosulfide phases. The observed difference suggests that the presence of SRB enables the transformation of polyphasic precursors to more crystalline structures through the combined effects of bacterial metabolites and localized precipitation within a low pH micro-environment around the cell walls. The addition of Fe(II) (i.e., $[Ni]_{a\alpha}/[Fe]_{a\alpha} = 5:1$) leads to formation of less crystalline Ni-sulfides in both biotic and abiotic systems, indicating crystal structure distortion caused by substitution of Ni with Fe. With decreasing [Ni]_{ao}/[Fe]_{ag}, Ni-sulfides become rarer, mixed Ni-Fe phases start to appear, and finally Ni-rich mackinawite (FeS) becomes the primary Ni-hosting phase at the lowest ratio tested ($[Ni]_{aq}/[Fe]_{aq} = 1:5$). We propose that whether aqueous Ni forms discrete Ni-S phases or is incorporated into dominantly

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126 INTRODUCTION

Nickel (Ni) is a bio-essential trace metal that is required for 128 129 several geochemically relevant metabolisms. These include the 130 cellular defense mechanism against reactive oxygen species, the production of ammonia from urea, the interconversion 131 of di-hydrogen and protons, and the microbially mediated 132 production of greenhouse gasses such as methane and carbon 133 monoxide (Ragsdale, 2009). Nickel concentrations are low in 134 135 the modern ocean with a range from 2 to 12 nM, and aqueous Ni shows a nutrient-like distribution with water depth 136 due to depletion in the photic zone as a result of biological 137 uptake (Sclater et al., 1976; Bruland et al., 2013). However, 138 elevated concentrations up to 2 mM can be found in streams 139 draining nickel-bearing minerals such as pyrite (FeS₂) and 140 millerite (trigonal NiS), as well as streams receiving industrial 141 and mining wastes (reviewed in Rinklebe and Shaheen, 2017). 142 High concentrations of Ni can adversely affect human health 143 and the ecosystem; therefore, the World Health Organization 144 recommended an upper limit of 0.001 mM (1 µM) for Ni 145 146 concentrations in drinking water (WHO, 2007). Consequently, 147 the removal mechanism of aqueous Ni from natural waters is of great interest both in terms of improving understanding 148 of natural biogeochemical processes and in applying that 149 information to improve water treatment quality. 150

sulfidic environments.

environments, mackinawite

In oxygen-rich environment, aqueous Ni in the water column 151 is initially scavenged by Fe-Mn-oxides and organic matter and 152 subsequently transported to the sediments (Sclater et al., 1976; 153 Gall et al., 2013). Development of anoxic conditions below the 154 sediment-water interface leads to the reduction of Fe-Mn-oxides 155 and the oxidation of organic matter, prompting the release of 156 Ni back into solution (Tribovillard et al., 2006; Weber et al., 157 2009; Hindersmann and Mansfeldt, 2014). In euxinic (i.e., anoxic 158 and sulfidic) zones generated by the metabolism of sulfate-159 reducing microbes, the released Ni can be re-sequestered either 160 161 by co-precipitation with Fe-sulfides such as mackinawite (FeS) 162 and pyrite (FeS₂) (Huerta-Diaz and Morse, 1992; Abraitis et al., 163 2004; Algeo and Maynard, 2004; Noël et al., 2015; Houben 164 et al., 2017; Ikogou et al., 2017) or by precipitation as discrete Ni-sulfides (Ferris et al., 1987). Between these two mechanisms, 165 co-precipitation with Fe-sulfides is typically the more important 166 removal mechanism due to the high abundance of Fe (~5 wt% 167 Fe vs. 0.005 wt% Ni in Earth's upper crust; Rudnick and Gao, 168 169 2003) and the slower water exchange kinetics of Ni compared to Fe, leading to preferential Ni incorporation into the faster-170 precipitating Fe-sulfides (Morse and Luther, 1999). 171

180 181 182 Discrete Ni-sulfides are rarely observed in nature (see 183 Huang et al., 2010 for a brief review of their occurrences). 184 Nonetheless, Ni-sulfides readily form in experiments employing 185 high concentrations of Ni comparable to those found in polluted 186 streams (Lewis and Swartbooi, 2006; Karbanee et al., 2008; 187 Cao et al., 2009; Sampaio et al., 2010; Reis et al., 2013; Kiran 188 et al., 2015). Additionally, geochemical observations indicate that 189 Ni removal is decoupled from Fe precipitation in sulfidic and 190 weakly acidic (pH 4-6) pore waters, suggesting the formation 191 of discrete Ni-sulfides under these conditions (Huerta-Diaz 192 et al., 1998; Frohne et al., 2011). Therefore, depending on the 193 environment, Ni-sulfides may be an important host for Ni. 194 The question of whether Ni is hosted in discrete Ni-sulfides or 195

within Fe-sulfides, and the relative importance of each, matters

because each phase has different solubilities and reactivities,

which subsequently governs the fate and bio-availability of Ni in

Fe-S phases is primarily determined by the precipitation kinetics, and our experiments

at [Ni]_{aq}/[Fe]_{aq} = 1:1 suggest that Ni-sulfide precipitation kinetics is comparable or

higher than Fe-sulfides at this condition. Overall, our study allows for prediction

on the phases and biogeochemical factors controlling Ni removal and availability in

Keywords: nickel sulfides, millerite, polydymite, vaesite, sulfate-reducing bacteria, polyphasic, sulfidic

nature (Thoenen, 1999). 199 Fortunately, the understanding on the low temperature 200 formation of Ni-sulfides has developed considerably over the 201 past 20 years. Most of the efforts have been focused on 202 abiotic precipitation with Ni as the sole metal species, yielding 203 Ni-sulfides such as hydrated NiS (Ni_{1.1} · 1.5 H₂O), millerite, 204 α-NiS, polydymite (Ni₃S₄), heazlewoodite (Ni₃S₂), and vaesite 205 (NiS₂) (Wang et al., 1997; Jeong and Manthiram, 2001; Huang 206 et al., 2009; Wilkin and Rogers, 2010; Table 1). The formed 207 Ni-sulfides are almost exclusively nanocrystalline. While these 208 past experiments provide a foundation for studies on Ni-sulfides, 209 they are not representative of natural systems where precipitation 210 occurs in complex solutions containing other metals and 211 microbial life. Specifically, the effects of both aqueous Fe(II) 212 and sulfate-reducing bacteria (SRB) (both of which should be 213 present in euxinic zones) on Ni-sulfide precipitation are relatively 214 understudied. The presence of Fe(II) during copper precipitation 215 for example has been shown to modify the mineralogy and 216 increase the crystallinity of copper-sulfides (Mansor et al., 2019). 217 The presence of SRB also modifies the formation mechanism, 218 size, shape, phase, and aggregation state of various metal sulfides 219 (Gramp et al., 2006; Moreau et al., 2007; Xu et al., 2016; Picard 220 et al., 2018). Specific to Ni, Sitte et al. (2013) found that crystalline 221 α -NiS formation was favored in cultures of SRB compared to 222 abiotic controls in which only amorphous NiS was precipitated. 223 The difference in the crystallinity between biogenic and abiogenic 224 precipitates can have important implications to the availability of 225 Ni in euxinic environments. 226

The current study is motivated by the need to identify 227 and characterize the Ni-hosting phases that can form in 228

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TABLE 1 Crystal chemistry for Ni-sulfide and Ni-Fe-sulfide minerals 220

	Crystal	Chemical	S/Me ratio	
Minerals	structure	composition		
Ni-S				
Hydrated NiS	-	Ni _{1.1} S · 1.5 H ₂ O	0.91	
Millerite	Trigonal	NiS	1.00	
α-NiS	Hexagonal	NiS	1.00	
Godlevskite	Orthorhombic	Ni ₉ S ₈	0.89	
Heazlewoodite	Trigonal	Ni ₃ S ₂	0.67	
Polydymite	Cubic	Ni ₃ S ₄	1.33	
Vaesite	Cubic	NiS ₂	2.00	
Ni-Fe-S				
Ni-rich mackinawite ^a	Tetragonal	$(Ni_xFe_{1-x})S; x \le 0.56$	1.00	
Pentlandite	Cubic	(Ni, Fe) ₉ S ₈	0.89	
Violarite ^b	Cubic	(Ni, Fe) ₃ S ₄	1.33	
Bravoite ^c	Cubic	(Ni, Fe)S ₂	2.00	

246 ^aBased on Wilkin and Beak (2017). ^bViolarite is a solid-solution between the 247 end-members polydymite and greigite (Fe₃S₄). It is most stable as FeNi₂S₄, with subsequent additions of Fe resulting in less stable phases (Vaughan and 248 Craig, 1985; Haider et al., 2012). ^cBravoite was discredited as a mineral by the 249 International Mineralogical Association. 250

natural euxinic environments. Correspondingly, we performed 252 a systematic set of experiments that mimic the biogeochemical 253 diversity of natural environments, where Ni is precipitated with 254 excess sulfide in both Fe-poor and Fe-rich systems [from "No Fe" 255 to aqueous Ni/Fe ratios ([Ni]_{aq}/[Fe]_{aq}) of 1:5 to 5:1] and in the 256 presence versus absence of SRB at near-neutral pH. We conclude 257 that across the [Ni]_{aq}/[Fe]_{aq} ranges typically observed in natural 258 environments, the primary Ni-hosting phase(s) in Fe-poor 259 systems will be Ni-sulfides (such as polyphasic precursors, α -NiS 260 and vaesite) while in Fe-rich systems, the primary Ni-hosting 261 phase will be mackinawite. The presence of SRB cells promotes 262 the growth and transformation of fine Ni-sulfide nanoparticles 263 into more crystalline phases, thus likely reducing the probability 264 of Ni re-mobilization into solution. 265

MATERIALS AND METHODS

All experimental procedures were performed under anoxic 270 conditions and using acid-cleaned containers unless otherwise 271 specified. Anoxic solutions were prepared by boiling for 20 min 272 and degassing with N₂ gas for another 20 min, followed by storage 273 under pure N2 or 97% N2 - 3% H2 headspace. The syntheses 274 were performed in a modified metal toxicity medium (MTM) 275 that contains (per liter): 3.8 ml 80% lactic acid, 2.23 g Na₂SO₄, 276 277 0.06 g CaCl₂.2H₂O, 1.0 g NH₄Cl, 1.0 g MgSO₄.7H₂O, 0.05 g yeast 278 extract, 0.5 g tryptone and 2.83 g HEPES. Initial pH was adjusted to 7.2 by the addition of NaOH. The initial concentrations of Ni 279 and Fe were adjusted by amending base MTM with metal stock 280 solutions of either 20 mM NiCl₂.6H₂O or 10 mM FeSO₄.7H₂O 281 (pH adjusted to 2 with HCl to minimize oxidation). The total 282 283 metal concentration was kept constant at 500 μ M, while the initial aqueous Ni/Fe ratios ([Ni]_{aq}/[Fe]_{aq}) were varied from 284 "No Fe," 5:1, 1:1, and 1:5. 285

Biotic Precipitation

Biotic syntheses were conducted in the presence of Desulfovibrio 287 vulgaris strain DSM 644 (also known as ATCC 29579 or strain 288 289 Hildenborough), which gradually produce sulfide as the end-290 product of sulfate reduction. The culture was initially maintained 291 over a period of months at 4°C in modified Postgate medium #63 292 (media composition detailed in Xu et al., 2016). Prior to each 293 round of synthesis, the culture was inoculated at 1% v/v into MTM and allowed to grow at 30°C while shaking. After 72 h, 294 the cells were inoculated at 1% v/v into 100 ml of Ni- and/or 295 296 Fe-amended MTM for the actual synthesis experiments. After 297 6-25 days of incubation, the culture and associated mineral 298 precipitates were aliquoted into two separate portions for 299 X-ray diffraction (XRD; ~80 ml) and transmission electron 300 microscopy (TEM; ~5 ml) analyses, respectively. Both aliquots 301 were centrifuged at 15,000 g for 5 min followed by removal of the supernatant via pipetting. Mineral pellets from the XRD aliquot 302 303 were rinsed once with 10 ml H₂O, resuspended in 0.8 ml H₂O 304 and then allowed to dry on a glass slide as a thin film. Mineral pellets for TEM were treated for 1 h at 45°C in a 10 ml solution 305 306 containing 0.1% Triton-X and 100 µg/ml each of lysozyme and 307 proteinase K to lyse the bacterial cells. Soluble organics were then 308 separated from the mineral fraction by centrifugation at 15,000 g for 15 min. The mineral pellets were rinsed twice more with 310 10 ml H₂O to remove traces of organics before being set aside for TEM analyses. 311

Abiotic Precipitation

For abiotic synthesis, 50 ml of 15 mM Na₂S.9H₂O solution was titrated at a rate of 2 ml/h into 50 ml of Ni- and/or Fe-amended MTM. The initial concentration of chemical components in the MTM were doubled so that after Na2S addition, the final concentration was similar as in the biotic synthesis. A Teflon-319 coated magnetic stir bar was used to mix the solution during 320 titration. After 5-36 days of incubation, nanoparticle aggregates 321 were separated from solution by filtration through a 0.2 μ m pore 322 size mixed cellulose ester membrane. The solution remained clear 323 during this step. The precipitates were then resuspended in 1 ml 324 H_2O and aliquoted for XRD (~0.8 ml dried onto glass slide) and 325 TEM (\sim 0.2 ml stored in sealed tubes) analysis. 326

Characterization

The samples were characterized using a combination of XRD 329 and TEM-based methods. The XRD patterns were collected using 330 a Rigaku Miniflex II from 10 to 60° 20 angle with a step size 331 of 0.05° and a scan speed of $0.5^\circ/\text{min},$ totaling a collection 332 time of about 1 h and 45 min per sample. Samples for TEM 333 were diluted in H₂O in the anaerobic chamber and shipped 334 in rubber-capped and aluminum-sealed glass serum vials to 335 the NanoEarth National Center at Virginia Tech. On site, the 336 samples were sonicated for up to 1 h to disperse aggregates prior 337 to mounting onto gold-coated TEM grids (product #01824G, 338 Ted Pella, Inc.). Microscopy images and selected area electron 339 diffraction patterns (SADP) were obtained using a JEOL-2100 340 TEM operating at 200 kV, and elemental ratios [e.g., S/(MeFe+Ni) 341 and Ni/Fe] were obtained with the accessory energy-dispersive 342

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TABLE 2 | Solution pH, solid-phase molar elemental ratios and identity of mineral phases (determined by XRD and SADP) in the samples. 343

Sample	Final pH	Description	Ni/Fe ratios ^a	S/Me ratios ^a	n	Phases ^c
Abiotic						
No Fe 8.23	Day 5	-	1.26 ± 0.10	5	pns, α-NiS ^d , pd ^d	
	Day 36	-	1.37 ± 0.05	2		
Ni/Fe = 5:1 8.24	Day 6	6.41 ± 1.34	1.41 ± 0.04	3	pns, pd ^d	
		Day 22	5.53 ± 1.22	1.15 ± 0.09	3	
Ni/Fe = 1:1 8.19	Day 30	1.01 ± 0.21	0.86 ± 0.12	3	pns, FeS/pn	
		Day 30 – spheroids ^b	1.63	0.78	1	
Ni/Fe = 1:5	8.03-8.11	Day 22	0.23 ± 0.04	1.15 ± 0.09	3	FeS, go/pn ^d
Biotic						
No Fe	7.12-7.25	Day 6	-	1.33 ± 0.03	8	pns, v ^d , pd ^d
		Day 22	-	1.81 ± 0.77	3	
Ni/Fe = 5:1	7.10–7.16	Day 6	10.56 ± 2.07	1.18 ± 0.06	6	pns, α-NiS ^d , pd ^d
		Day 12	4.09 ± 1.18	1.95 ± 0.64	3	
Ni/Fe = 1:1	7.18	Day 6	1.11 ± 0.04	1.00 ± 0.01	2	pns, FeS, vi/pd
		Day 6 – spheroids ^b	3.84 ± 0.04	0.97 ± 0.01	3	
		Day 25	1.00 ± 0.30	1.09 ± 0.07	3	
Ni/Fe = 1:5	7.04-7.07	Day 6	0.20 ± 0.04	1.15 ± 0.04	4	FeS, pns ^d
		Day 6 – spheroids ^b	1.24 ± 0.89	1.10 ± 0.07	8	

363 ^aMeasured by EDS; ^bArea-targeted EDS spot analyses on spheroid-rich regions. Some Fe-rich nanosheets are also sampled due to their tight association with the 364 spheroids; ^cpns, polyphasic Ni-sulfide precursors; pd, polydymite; v, vaesite; vi, violarite; go, godlevskite; pn, pentlandite; ^dMinor phases.

X-ray spectroscopy (EDS) system coupled to a silicon drift 366 367 detector. Spot EDS analysis used beam sizes of 20-500 nm in diameter, while mapping analysis was conducted in scanning-368 369 TEM mode with a resolution of ~ 10 nm. The inherent error 370 of $\leq 7\%$ for the standardless EDS analysis method (Newbury 371 and Ritchie, 2014), compounded by the possibility of sulfur 372 adsorption onto the nanoparticles' surface (Rickard et al., 2006; Huang et al., 2010; Picard et al., 2016) may contribute to 373 374 the deviation of the measured elemental ratios from the real 375 ones, especially for nanoparticles with substantially high surface areas. Thus, the EDS data were considered semi-quantitative. 376 There were no observable differences in the nanoparticles' 377 characteristics between the samples that were dried under anoxic 378 379 conditions versus in air (<10 min drying time). There were also no obvious enhanced dissolution (e.g., less yield over time, 380 381 etch pits, noticeably different morphology) for the samples that 382 were stored in H₂O for up to 2 months prior to mounting 383 versus samples that were mounted to TEM grids immediately following preparation. Some Fe-rich samples that were stored in 384 vacuum for >1 week after mounting onto TEM grids did show 385 386 morphological changes that are likely caused by oxidation to Fe-387 oxides. All the samples in this study were therefore analyzed 388 within 24 h of mounting onto TEM grids (while kept within a dry 389 vacuum chamber during this period). Particle size distributions 390 were obtained by measuring the diameter of >50 individual 391 nanoparticles manually in ImageJ¹.

RESULTS

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Within a day of the start of incubation, black or dark brownish 396 precipitates are readily formed under all tested conditions. Starting from an initial pH of 7.2, the final pH values of the biotic experiments decrease slightly to 7.15 ± 0.10 . In contrast, the final solution pH values in all of the abiotic experiments are about one unit higher, with a range of 8.03-8.24 (Table 2).

Characterization of Nanoparticles Precipitated in Fe-Free and Fe-Poor Systems

Analyses of XRD patterns of precipitates formed in the 432 Fe-free and Fe-poor abiotic and biotic systems ("No Fe" and 433 $[Ni]_{aq}/[Fe]_{aq} = 5:1$ indicate that the precipitates are largely 434 amorphous (Figure 1). Microscopy observations of early-stage 435 (<1 week old) samples reveal a predominance of Ni-sulfide 436 spheroids (Figure 2), with particle sizes typically <10 nm 437 and extending up to ~ 20 nm (Figure 3). The SADP derived 438 from these nanoparticles reveal two major diffuse rings with 439 d-spacing ranges of 2.44-2.83 Å and 1.64-1.99 Å (Figure 2c). 440 The d-spacings are consistent with those of millerite, although 441 we cannot rule out the presence of other phases such as α -NiS, 442 godlevskite and polydymite (Supplementary Table S1). The 443 average sulfur-to-metal ratios for these spheroids are slightly 444 above 1.00 (Table 2), which is likely due to excess sulfur 445 adsorption to nanoparticles' surfaces (Rickard et al., 2006; 446 Huang et al., 2010; Picard et al., 2016). High-resolution TEM 447 images reveal the polyphasic nature (as defined by French 448 et al., 2012) of the early-stage Ni-sulfide spheroids, in that they 449 contain crystalline areas \sim 1–3 nm in size within an amorphous 450 matrix (Figure 2b). From here on, we will refer to these 451 spheroidal nanoparticles as "polyphasic Ni-sulfide precursors." 452 With increasing incubation time, the polyphasic precursors 453 display enhanced crystallinity and a distribution toward larger 454 sizes (Figures 2d,e, 3). The biotic precipitates in particular are 455 generally more crystalline and larger than abiogenic precipitates. 456

¹https://imagej.nih.gov/ij/ 399

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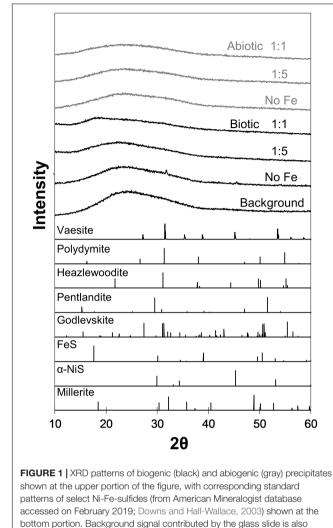
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included in the figure. Most samples are indistinguishable from the background except "Biotic No Fe" and "Biotic 1:1 Ni:Fe" samples.

Distinct diffraction spots become noticeable in the SADP of the aged or biogenic crystalline nanoparticles (**Figure 2f**). The measured d-spacings from the spots are 2.91, 2.82, 2.54, 2.40–2.43, 2.21, 1.90, and 1.70 Å, which can be best explained by the reflections from millerite's (101), (300), (021), (220), (211), (131), and (401) planes, respectively.

Other Ni-sulfide phases such as vaesite, α -NiS and polydymite/violarite are also positively identified in the samples. Vaesite occurs as \sim 20-nm spheroidal nanoparticles or 100- to 300-nm euhedral nanocrystals only in the biotic system and in the absence of Fe (Figures 4a,b). The presence of vaesite is confirmed via both XRD (small peaks at 31.5°, 41.5°, and 54° 508 2θ in **Figure 1**) and SADP (insets of **Figures 4a,b**) analyses. The α -NiS phase occurs as \sim 20- to 80-nm hexagonal nanoplates and 509 is detected in aged (>1 week old) abiogenic precipitates in the 510 511 absence of Fe and in biogenic precipitates at $[Ni]_{aq}/[Fe]_{aq} = 5:1$ (Figure 4c). Finally, polydymite (Ni₃S₄)/violarite (FeNi₂S₄) 512 occurs as ~10-20 nm irregularly shaped nanoparticles in all of 513

the Ni-rich samples (**Figure 4d**). This phase is identified based 514 on high-resolution TEM images that reveal lattice fringes with 515 3.32–3.38 Å d-spacing corresponding to the (220) planes of 516 polydymite/violarite (**Supplementary Table S1**). 517

Characterization of Nanoparticles Precipitated in Fe-Rich Systems

Analysis of TEM and EDS data of biotic and abiotic samples 522 precipitated at [Ni]_{aq}/[Fe]_{aq} = 1:1 indicate the presence of Ni-523 rich spheroids (Ni/Fe ratio up to 3.84) that are tightly associated 524 with extensive Fe-Ni-sulfide nanosheets (Ni/Fe = 1.01-1.11) 525 (Figures 5a,b, 6). Compared to those formed in the Fe-poor 526 systems, the Ni-rich spheroidal nanoparticles here have lower 527 crystallinity and are finer in size (Figure 3). The mineralogy 528 of the spheroids and nanosheets are complex. Morphological 529 similarities suggest that the Ni-rich spheroids are polyphasic 530 Ni-sulfide precursors whereas the Fe-Ni-sulfide nanosheets are 531 Ni-containing mackinawite (see Mansor et al., 2019 for images 532 of Ni-free mackinawite formed under comparable conditions). 533 Based on the XRD patterns, these precipitates are largely 534 amorphous except that a tentative shoulder occurs at 17° 20 535 (\sim 5 Å), attributable to the (001) planes of mackinawite, in 536 the biogenic samples (Figure 1). In the abiogenic precipitates, 537 SADP collected from spheroid-rich regions indicate a diffuse 538 ring with a d-spacing range of 2.40-2.82 Å, and additional 539 diffraction spots with d-spacings of 2.51, 2.43, and 1.70 Å 540 (Figure 5a, inset); these d-spacings are most consistent with those 541 of millerite. Comparatively in the biogenic precipitates, SADP 542 collected from spheroid-rich regions display diffraction rings 543 with measured d-spacings of 3.38, 2.85, 2.42, 1.84, and 1.71 Å 544 (Figure 5b, inset) that match the reflections from major planes 545 of violarite/polydymite. Additional d-spacings obtained from 546 imaging of lattice fringes are: (a) 5.4–5.7 Å from the nanosheets, 547 most likely corresponding to reflections from the (001) plane 548 of disordered mackinawite (Wolthers et al., 2003) while not 549 ruling out overlapping reflections from pentlandite (111) and/or 550 violarite (111) planes, and (b) \sim 1.5 Å from the Ni-rich spheroids, 551 corresponding to the (012) planes of millerite. Overall, the 552 precipitates are likely mixtures of the aforementioned phases with 553 varying degrees of crystallinity at various combination ratios. 554

As the initial relative $[Fe]_{aq}$ is increased to a $[Ni]_{aq}/[Fe]_{aq}$ 555 ratio of 1:5, the precipitates become dominated by Fe-rich 556 nanosheets (with a Ni/Fe ratio of ~ 0.20) (Figures 5c,d). 557 Analyses of these samples' XRD patterns confirm the presence of 558 mackinawite (Figure 7). The SADP analyses are also consistent 559 with the reflections of mackinawite especially for the biogenic 560 precipitates (Figures 5c,d, insets). For the abiogenic precipitates, 561 additional rings with corresponding d-spacings of 3.57 and 562 2.89 Å indicate the presence of pentlandite or godlevskite along 563 with mackinawite. Interestingly, while no discrete Ni-sulfides are 564 observed in the abiogenic precipitates, the biogenic precipitates 565 still contain a significant number of Ni-rich spheroids (Ni/Fe 566 ratio = 1.24 ± 0.89 ; n = 8). These spheroidal nanoparticles have 567 a mean size of 11.9 ± 3.2 nm (n = 8). Therefore, the presence of 568 D. vulgaris plays a role in favoring the precipitation of discrete 569 Ni-sulfides even in Fe-rich systems. 570

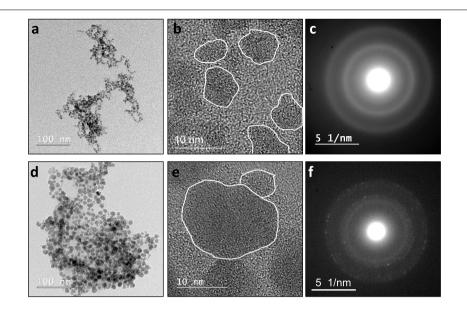


FIGURE 2 Left to right: low resolution images, high resolution images and SADPs derived from the spheroidal nanoparticles. **Top row (a–c)**: spheroidal nanoparticles observed in abiotic "No Fe" systems during the early periods of incubation (\leq 1 week old). These nanoparticles are polyphasic, in that they display crystalline nano-domains ~1–3 nm in size within an amorphous matrix. The corresponding SADP displays only diffuse rings. **Bottom row (d–f)**: spheroidal nanoparticles typically observed in the biotic "No Fe" systems. These nanoparticles are larger in size and are more crystalline. Similar nanoparticles are also observed in the abiotic "No Fe" systems. The corresponding SADP starts to resolve into distinct spots with a set of d-spacing values that are most consistent with those of millerite. Annotated SADPs are available in **Supplementary Material**.

DISCUSSION

Formation Mechanism of Ni-Sulfides

Through the various techniques employed in our study, we have identified several phases of Ni-sulfides including polyphasic Nisulfide precursors, millerite, α -NiS, polydymite and vaesite. Each of these phases has different solubility and formation energy, with vaesite being predicted as the thermodynamically most stable phase at room temperature (Wilkin and Rogers, 2010). Since all of the Ni-sulfides are precipitated as nanoparticles, their solubilities are also affected by their sizes and crystallinity (Gilbert and Banfield, 2005; Hochella et al., 2008; Caraballo et al., 2015). Understanding the factors that control the formation of a specific Ni-sulfide phase, as well as their size and crystallinity, can help us better predict the behavior and fate of Ni in euxinic environments. Below, we discuss the formation pathway(s) of Ni-sulfides in the absence and presence of SRB, and how these pathways may be modified in the presence of Fe(II).

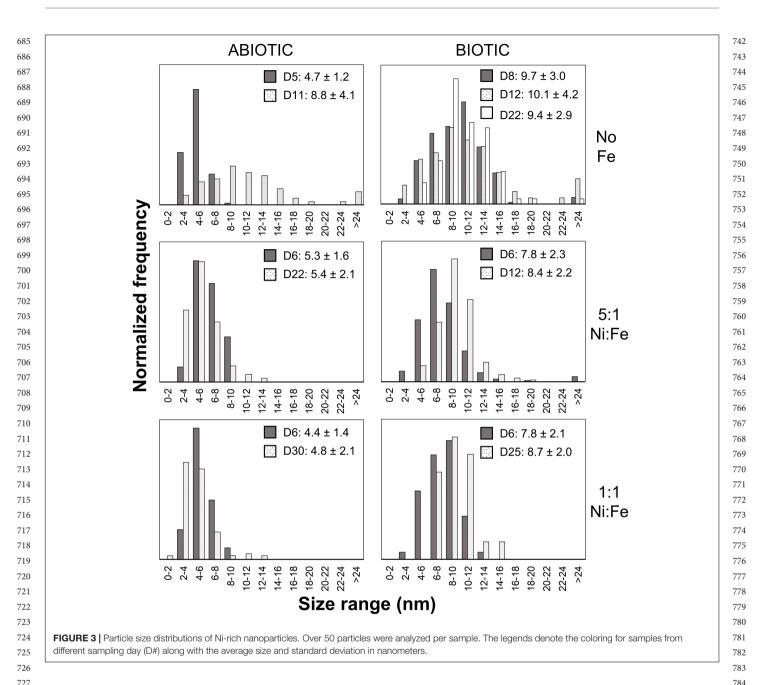
617 Abiotic Formation of Ni-Sulfides

In our study, the first Ni-sulfide phase to precipitate out of solution is XRD-invisible polyphasic Ni-sulfide precursors. Previous abiotic studies indicate that the first product between aqueous Ni and sulfide at 25°C is also XRD-invisible Ni-sulfides (Wang et al., 1997; Jeong and Manthiram, 2001; Wilkin and Rogers, 2010). This phase is later determined through a combined TEM and pair distribution function analysis to be nanoparticulate Ni-sulfides with a core-shell structure (Huang et al., 2009, 2010). The 1-3 nm core is composed of crystalline millerite while the 3-5 nm thick shell is

composed of slightly Ni-rich material with structural H_2O in an amorphous configuration, leading to the chemical composition of Ni_{1.1}S · 1.5H₂O. Overall, the structure and size of this hydrated phase is similar to the polyphasic Ni-sulfide precursors identified in our study. 659

Over time in the abiotic systems, we observe the development of early-stage polyphasic Ni-sulfide precursors to larger and more crystalline nanoparticles up to α -NiS within a few days to weeks. Wilkin and Rogers (2010) observed the formation of α -NiS within similar timescales at room temperature, and its formation rate is accelerated to within a day at 60°C. The presence of stacking faults and twinning within the nanoparticles in our study (Figure 8) implicate crystal growth mechanism through particle attachment of smaller nanoparticles or soluble nanoclusters, with potential subsequent atomic rearrangements to higher stability structures (De Yoreo et al., 2015). This is consistent with solubility consideration; since most metal sulfides have low solubilities, the degree of supersaturation during precipitation is extremely high (Rickard and Luther, 2006). Initial mineral formation is therefore dominated by nucleation events, and additional growth can only proceed through attachment and re-assembly of the initially formed nanoparticles. Hexagonal α-NiS is generally considered as a metastable phase but once formed, it is stable for over 6 months in aqueous solution at room temperature (Wang et al., 2006; Wilkin and Rogers, 2010).

Careful examination of the abiogenic precipitates in our study $_{680}$ also indicates the presence of trace amounts of polydymite, $_{681}$ but not vaesite in the Fe-poor systems. Wilkin and Rogers $_{(2010)}$ have shown that in the presence of elemental sulfur, $_{683}$ $_{\alpha}$ -NiS is rapidly transformed to vaesite through polydymite as $_{684}$



an intermediate (**Figure 9**). The added elemental sulfur acts as an oxidant to facilitate the partial oxidation of Ni (from Ni²⁺ in α -NiS to Ni³⁺ in polydymite) and to produce polysulfides (e.g., S₂²⁻ in vaesite) by reacting with H₂S. Thus, the pathway of vaesite formation is comparable to one of the pathways to forming pyrite in the Fe-S system, in which mackinawite is transformed to pyrite via greigite intermediate (Wang and Morse, 1996; Benning et al., 2000; Hunger and Benning, 2007):

XRD-amorphous NiS $\rightarrow \alpha$ -NiS \rightarrow polydymite (Ni₃S₄)

 \rightarrow vaesite (NiS₂)

 $\begin{array}{cc} ^{739} \\ 740 \end{array} \qquad \qquad \text{XRD-amorphous FeS} \rightarrow \text{mackinawite (FeS)} \end{array}$

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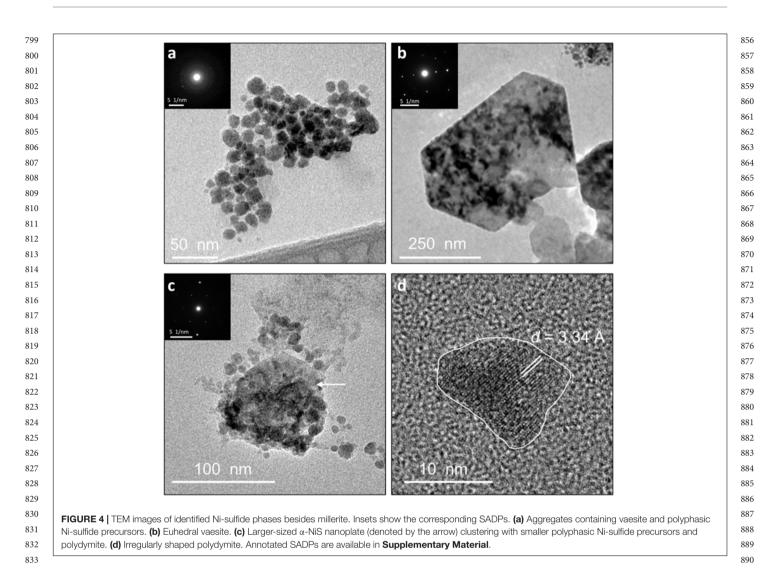
 \rightarrow greigite (Fe₃S₄) \rightarrow pyrite (FeS₂)

The formation of polydymite and vaesite are favored at acidic 785 pH (Wang et al., 1997; Huang et al., 2009). By analogy to the 786 Fe-S system, this preference is attributed to (i) protons acting as 787 oxidants for the transformation of MeS to Me₃S₄ (Bourdoiseau 788 et al., 2011) and (ii) higher rate of dissolution-reprecipitation 789 reactions to form MeS₂ from either MeS or Me₃S₄ (Schoonen 790 and Barnes, 1991; Wang and Morse, 1996). The lack of vaesite 791 in the abiotic systems in our study is consistent with predictions 792 for the near-neutral pH experimental conditions and suggests 793 that negligible amounts of polysulfides are formed (these species 794 may form in the presence of trace oxygen). In contrast, the trace 795 amount of polydymite found in all the Fe-poor abiotic samples 796 is rather surprising. Previous understanding suggests that this 797 phase is primarily an artifact of sample preparation and is 798

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formed during the brief period of oxidation either during sample 835 mounting or during transport to the TEM rather than being 836 an actual experimental product. This is similar to the multiple 837 reports of greigite detection in experimental Fe-S systems, which 838 has been attributed to the fast oxidation of mackinawite to 839 greigite in air (Boursiquot et al., 2001) and even under vacuum 840 (reviewed by Rickard and Luther, 2007). Indeed, solid-state 841 oxidation of hydrated NiS to polydymite has been shown to occur 842 even under an anoxic atmosphere with purportedly <2 ppm O₂ 843 (Huang et al., 2009). 844

846 Formation of More Crystalline Biogenic Ni-Sulfides

Precipitation in the presence of D. vulgaris leads to more 847 848 crystalline structures of Ni-sulfides as indicated by two lines of observations. First, higher stability phases such as vaesite (in "no 849 Fe" systems) and polydymite/violarite (at [Ni]_{aq}/[Fe]_{aq} = 1:1) 850 are more prominent in the biotic compared to abiotic 851 852 systems. Second, biogenic Ni-sulfide nanoparticles are larger in size and more crystalline than their abiogenic counterparts 853 precipitated at comparable [Ni]_{aq}/[Fe]_{aq} and incubation time. 854 These observations imply a faster growth and transformation rate 855

of initially formed Ni-sulfides to more crystalline structures when precipitation occurs in the presence of SRB.

We first consider if the slightly lower pH in the biotic experiments (final pH \sim 7) compared to the abiotic experiments (final pH \sim 8) may increase the transformation rate of Nisulfides. The pH difference can be understood by considering the protons and buffering capacity generated during microbial sulfate reduction reaction (Pankhania et al., 1986; Voordouw, 2002; Heidelberg et al., 2004): 900

$$2 C_2 H_4 OHCOO^- + SO_4^{2-}$$
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$$\rightarrow 2 \text{ CH}_3 \text{COO}^- + 2 \text{ HCO}_3^- + \text{HS}^- + \text{H}^+$$
 (3)

compared to the protons consumed during the dissociation reaction of sodium sulfide in the abiotic systems:

$$Na_2S + H^+ \rightarrow Na^+ + HS^-$$
 (4) 907
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Although the lower pH in the biotic systems may contribute 909 to the more crystalline structures of biogenic Ni-sulfides by (i) 910 favoring the formation of polydymite and vaesite (see section 911 "Abiotic Formation of Ni-Sulfides") and (ii) favoring growth 912

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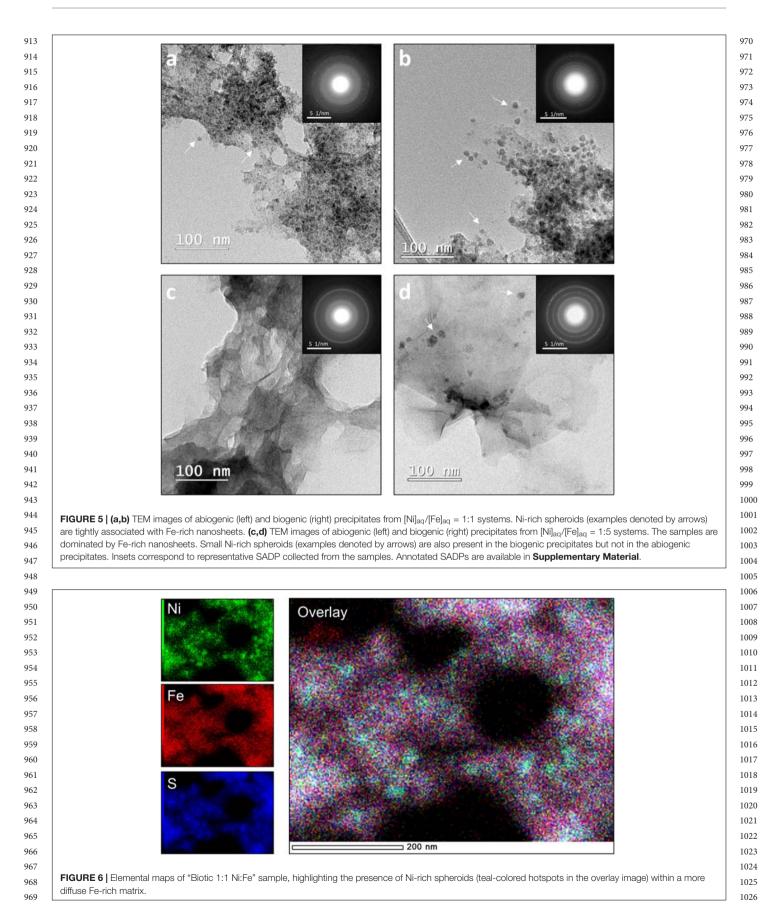
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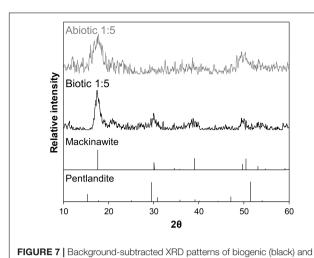
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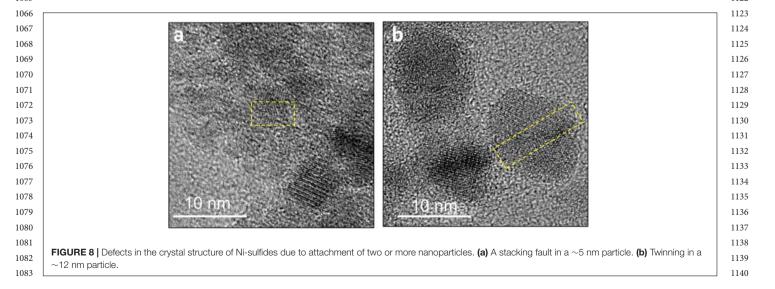


abiogenic (gray) precipitates at $[Ni]_{aq}/[Fe]_{aq} = 1:5$ shown at the upper portion of the figure, with corresponding standard patterns of mackinawite and pentlandite (from American Mineralogist database; Downs and Hall-Wallace, 2003) shown at the bottom portion. The sample patterns are most consistent with mackinawite.

to larger-sized Ni-sulfides by slowing down the nucleation rate 1049 (metal solubility tends to increase with decreasing pH; Lewis, 1050 2010), we argue that the observed pH difference is insufficient 1051 to cause these effects. First, previous abiotic experiments have 1052 shown that Ni-sulfides precipitated at room temperature from 1053 pH 3-9 share similar initial mineralogy (Huang et al., 2009). 1054 Second, the abiotic formation of polydymite and vaesite requires 1055 1056 a pH range that is much more acidic (pH \leq 5; Wang et al., 1997; 1057 Huang et al., 2009) than that measured for the biotic systems. 1058 Third, the most recent model of Ni-sulfide solubility in the 1059 presence of excess sulfide indicates that the solubility does not change appreciably across pH 7-8 (Figure 10). Considering all of 1060 these, we argue that the formation of more crystalline biogenic 1061 Ni-sulfides are not mainly caused by the slight variation in the 1062 bulk solution pH. Other effects of SRB to Ni-sulfide precipitation 1063 must be considered. 1064

Besides altering the inorganic chemistry of the solution, the 1084 presence and metabolism of SRB are also known to provide 1085 additional precipitation sites for metal sulfides. Binding of 1086 Ni to cell walls and extracellular polymeric substances (EPS) 1087 secreted by bacteria are well-known (Beveridge and Murray, 1088 1976; Beveridge and Koval, 1981; Ferris et al., 1989; Wuertz 1089 et al., 2000; Guibaud et al., 2005; Quintelas et al., 2009) and 1090 Ni-sulfides have been observed to precipitate on bacterial cell 1091 walls under both experimental (Fortin et al., 1994; Sitte et al., 1092 2013) and natural (Ferris et al., 1987) conditions. During the 1093 initial nucleation event, it is important to recognize that the 1094 (sub)micro-environments around the cell walls are chemically 1095 distinct from the bulk solution due to limitations in diffusion at 1096 small scales (Purcell, 1977). A prominent feature of this micro-1097 environment is a high concentration of sulfide, which creates 1098 local supersaturation zones that favor metal sulfide precipitation 1099 (Southam, 2000). Additionally, the protons and organics (e.g., 1100 acetate) produced as by-products of sulfate reduction (Eq. 3) will 1101 be concentrated in this micro-environment. In other bacterial 1102 cultures and natural biofilms, the micro-environment pH has 1103 been shown to reach levels as low as 3.5 compared to the near-1104 neutral bulk solution (Mera et al., 1992; De Los Ríos et al., 2003; 1105 Hunter and Beveridge, 2005; Hidalgo et al., 2009). A combination 1106 of micro-environmental low pH and the presence of soluble 1107 organics that can bind Ni are likely contributing to the distinct 1108 characteristics of biogenic Ni-sulfides. 1109

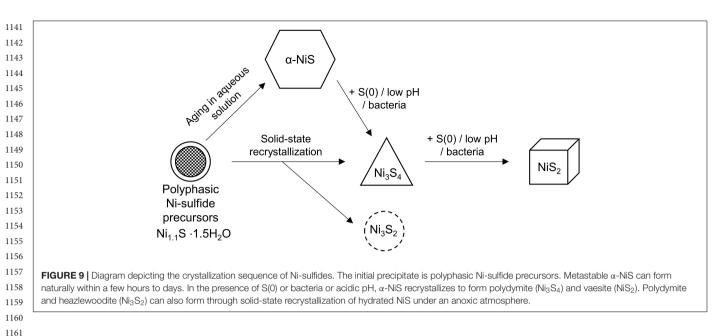
As previously mentioned, the formation of vaesite and 1110 polydymite are favored at acidic pH (pH <5), which plausibly 1111 exist in the micro-environments around SRB cells. The fact that 1112 neither of these phases are the most abundant phase in the biotic 1113 systems indicates some chemical/physical limitations to their 1114 formation, largely in line with the micro-environment formation 1115 hypothesis. Once formed, these phases may detach from the cell 1116 walls but are nevertheless stable at the near-neutral pH of the bulk 1117 solution (Wilkin and Rogers, 2010). Furthermore, precipitation 1118 in an acidic micro-environment may also explain the larger sizes 1119 of biogenic Ni-sulfides compared to their abiogenic counterparts. 1120 An analogy can be made using the example of mackinawite and 1121



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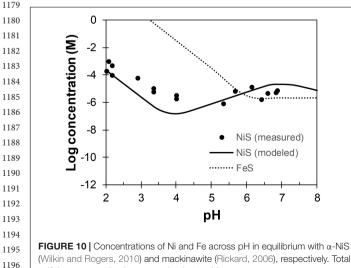
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pyrite, both of which tend to be larger in size when precipitated at 1162 acidic pH (Schoonen and Barnes, 1991; Wang and Morse, 1996). 1163 This phenomena is attributed to the increase in metal solubility 1164 with decreasing pH, leading to a decrease in the degree of 1165 supersaturation that results in slower nucleation rates and faster 1166 growth rates compared to the scenarios at higher pH (Lewis, 1167 2010). The solubility of Ni, however, changes in complex fashion 1168 across pH and only increases with decreasing pH below pH 4 in 1169 the presence of sulfide (Figure 10). Therefore, the formation of 1170 more crystalline Ni-sulfides likely requires a micro-environment 1171 pH \approx 4, which remains to be confirmed experimentally for SRB. 1172

In addition to pH modification, the organics secreted by SRB can also affect the dynamics of Ni-sulfide precipitation. For example, bacterial metabolites have been implicated in mediating attachments between ZnS nanoparticles, allowing biogenic ZnS to achieve larger sizes compared to abiogenic



sulfide concentration is assumed to be 1 mM.

ZnS (Xu et al., 2016). Such a mechanism may also operate in our study, but we note that there are no indications of more particle attachment events (e.g., defects such as stacking faults and twinning) in the biogenic Ni-sulfides compared to abiogenic Ni-sulfides. Secreted proteins may also decrease the availability of free Ni(II) ion by complexation (Fortin et al., 1994), resulting in lower degree of supersaturation and contributing to larger Ni-sulfides by altering the relative kinetics of Ni-sulfide growth over nucleation. Overall, we propose that the combination of precipitation within acidic micro-environments and the interaction of Ni/Ni-sulfides with secreted organics act in tandem to generate more crystalline Ni-sulfides in the biotic systems.

Metal Sulfide Formation During Co-existence of Fe and Ni: Effects of Cation Sizes and Reaction Kinetics

1233 In the presence of appreciable quantities of Fe(II)_{aq}, a diversity of 1234 Ni-hosting phases is formed including Ni-sulfides, mackinawite 1235 and (minor) mixed Ni-Fe phases such as pentlandite and 1236 violarite. Pentlandite is generally considered a high temperature 1237 phase of Fe-Ni-sulfides that weathers to form violarite at low 1238 temperatures (e.g., Grguric, 2002; Tenailleau et al., 2006; Xia et al., 1239 2009), but its stability field could be enhanced at the nano-scale 1240 due to changes in the relative stability of minerals as a function of 1241 particle size (Navrotsky et al., 2008). The preferential formation of 1242 all of the aforementioned phases and their respective properties 1243 are strongly influenced by the solution's initial [Ni]aq/[Fe]aq 1244 ratios and can be understood by considering the differences 1245 in the bond chemistry, reaction kinetics and potential post-1246 precipitation recrystallization of Ni- and Fe-sulfides. 1247

When the initial concentration of Ni exceeds Fe ($[Ni]_{aq}$ / 1248 [Fe]_{aq} = 5:1), our TEM analyses indicate that Fe is associated 1249 with Ni-sulfides rather than forming separate Fe-sulfide phases. 1250 This is consistent with the observation of the presence of Fe 1251 impurities in natural millerite and polydymite (Belkin and Luo, 1252 2008). It is apparent through our experiments that the addition 1253 of Fe(II) leads to the formation of smaller, less crystalline 1254

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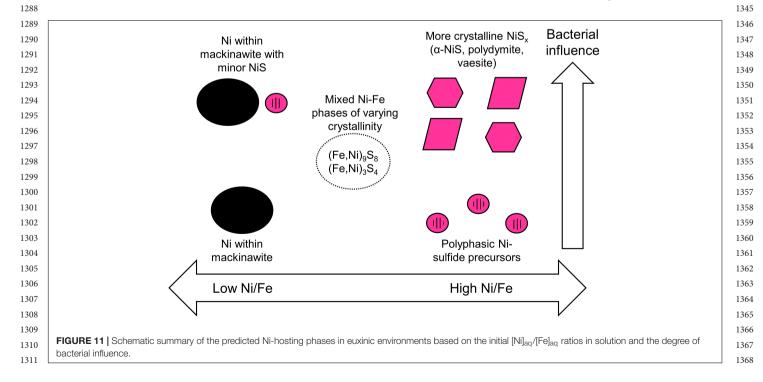
Ni-sulfide nanoparticles (Figure 3). We hypothesize that this 1255 can be explained if Fe is incorporated into the Ni-sulfide crystal 1256 lattice. Due to the \sim 9% larger ionic radius of Fe(II) than Ni(II) 1257 (Shannon, 1976), the incorporation of Fe into Ni-sulfides can lead 1258 to structural instability that prevents or delays the formation of 1259 more crystalline Ni-sulfides. Confirmation of this factor depends 1260 on a more comprehensive approach using density functional 1261 theory to account for complex bonding energetics of transition 1262 metals in sulfide minerals (Kwon et al., 2015). 1263

At equimolar concentrations of Ni and Fe ([Ni]aq/ 1264 $[Fe]_{aq} = 1:1$, two classes of minerals are appreciably present: 1265 Ni-sulfide spheroids (predominantly polyphasic Ni-sulfide 1266 precursors of millerite and minor polydmite/violarite) and 1267 Fe-rich nanosheets (predominantly mackinawite and likely some 1268 pentlandite based on the electron diffraction patterns). The 1269 1270 preferential partitioning of Ni and Fe into two separate phases is consistent with the notion that millerite and mackinawite are 1271 not likely to be miscible based on the differences in their S-Me 1272 bond distances and coordination numbers (Wilkin and Beak, 1273 2017). In the absence of significant post-precipitation mixing, 1274 1275 metal partitioning is therefore more likely to be governed by the relative early-stage precipitation kinetics of Ni- and Fe-1276 sulfides. The water exchange kinetics of Ni is about two orders 1277 of magnitude slower than Fe, prompting suggestions that Ni 1278 should preferably be incorporated into Fe-sulfides rather than 1279 forming discrete Ni-sulfides (Morse and Luther, 1999). However, 1280 the abundance of Ni-sulfides in our experiments indicates that 1281 Ni precipitates faster than (or at least comparable to) Fe at 1282 equimolar concentration. This implies that the simple kinetic 1283 approach taken previously can lead to erroneous predictions. 1284 Rather, a more sophisticated kinetic model that incorporates 1285 1286 other factors such as the degree of supersaturation, surface areas 1287 and activation energies need to be developed for metal sulfides.

Finally, at the lowest ratio tested ($[Ni]_{aq}/[Fe]_{aq} = 1:5$), 1312 mackinawite is the dominant phase in these systems. In the 1313 abiotic systems, Ni is fully associated with mackinawite either 1314 by adsorption or by incorporation to form a solid-solution 1315 (Fe_{1-x}Ni_xS where maximum x = 0.56; Wilkin and Beak, 2017). 1316 This observation implies that mackinawite forms more rapidly 1317 than Ni-sulfides in the abiotic systems at this concentration 1318 ratio. Interestingly though, discrete Ni-sulfides are still present 1319 in the biotic systems at this concentration ratio. This implies 1320 that the precipitation kinetics of Ni is enhanced in the 1321 presence of D. vulgaris, which is consistent with the micro-1322 environment precipitation theory. As previously discussed, Ni-1323 sulfide solubility decreases with lower pH between pH 4 and 1324 7 while Fe-sulfide solubility increases over the same pH range 1325 (Figure 10). This will cause an increase of Ni precipitation 1326 kinetics relative to Fe, thus allowing the formation of some 1327 discrete Ni-sulfides that will persist once formed. 1328

ENVIRONMENTAL IMPLICATIONS

This study explores the form and attributes of Ni-hosting 1333 phases in euxinic settings that can precipitate across a range 1334 of aqueous Ni-to-Fe ratios and in the presence and absence 1335 of SRB. The variables tested in this study are chosen so as to 1336 capture the biogeochemical diversity of potential Ni precipitation 1337 sites in natural settings. For example, SRB is known to be a 1338 primary producer of sulfide in low temperature environments 1339 and are expected to be intimately associated with metal sulfide 1340 precipitation (Picard et al., 2016). Experiments, however, have 1341 shown that even when SRB are present, a significant fraction of 1342 the metals may still be precipitated in the bulk solution away from 1343 the cell surfaces (Picard et al., 2018; Stanley and Southam, 2018). 1344



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Therefore, it is relevant to consider and expect that both abiotic 1369 and biotic precipitation will occur in tandem in nature. In 1370 terms of metal concentrations, we utilize high concentrations 1371 that are more similar to highly polluted streams rather than 1372 pristine environments. The chosen initial aqueous Ni-to-Fe ratios 1373 are comparable to those observed in nature, ranging from 1374 Ni-rich and Fe-poor streams receiving inputs from industrial 1375 wastewaters, mining wastes and Ni-rich bedrocks (e.g., Periasamy 1376 and Namasivayam, 1995; Heikkinen et al., 2002; Moreton et al., 1377 2009; Azizullah et al., 2011) to relatively Fe-rich and Ni-poor 1378 water bodies as observed primarily in pristine environments 1379 (oceans, rivers, groundwaters) and also in some polluted streams 1380 1381 (Thoenen, 1999; Lee et al., 2002; Rahman et al., 2015). The results of our study therefore allow predictive capabilities on the phases 1382 controlling Ni availability in euxinic environments across a range 1383 1384 of aqueous Ni/Fe ratios (Figure 11).

Based on our study, we will expect nanoparticulate Ni-1385 sulfides to be important Ni-hosting phases in Fe-poor euxinic 1386 environments. These Ni-sulfides will age and develop over time 1387 to more crystalline phases particularly in the presence of SRB. 1388 Trace amounts of Fe, however, will likely delay the development 1389 to more crystalline phases, making it more likely for Ni to be 1390 re-released into solution especially under conditions of rapidly 1391 alternating redox potential. 1392

As Fe concentration is increased relative to Ni, one needs to 1393 consider the relative fraction of Ni sequestered in Ni-sulfides 1394 versus those in association with Fe-sulfides. Our experimental 1395 results at $[Ni]_{aq}/[Fe]_{aq} = 1:1$ indicate that Ni-sulfides will be 1396 the primary Ni-hosting phase under this condition, with other 1397 Fe-sulfide phases being of secondary importance. When Fe 1398 concentration exceeds Ni, one should expect Fe-sulfides such as 1399 1400 mackinawite to be the primary Ni-hosting phase. Even though 1401 discrete Ni-sulfides are still precipitated in the biotic systems at $[Ni]_{aq}/[Fe]_{aq} = 1:5$, mass balance dictates that this phase 1402 contain at most 7% of the initial Ni fraction (assuming all 1403 metals precipitated, a maximum Ni/Fe ratio of 3.84 for the Ni-1404 sulfides and minimum Ni/Fe ratios of 0.16 for the nanosheets). 1405 It is important to note that Ni incorporation increases the 1406 growth kinetics and thermodynamic stability of mackinawite, 1407 with potential implications to the fate of Ni and other various 1408 trace metals in natural environments (Kwon et al., 2015; 1409 Ikogou et al., 2017). 1410

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DATA AVAILABILITY

All datasets generated for this study are included in the manuscript and/or the Supplementary Files. 1429

AUTHOR CONTRIBUTIONS

MM and JX designed the study. MM performed all the 1434 1435 experiments and wrote the first draft of the manuscript. MM 1436 and CW conducted the analyses. All authors contributed 1437 manuscript revision and read and approved the to 1438 submitted version.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: https://www.frontiersin.org/articles/10.3389/feart. 2019.00151/full#supplementary-material

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Conflict of Interest Statement: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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